

Hierarchical waveform inversion with double beamforming

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SUMMARY

Full waveform inversion is a promising imaging technology to delineate high resolution images of the subsurface, but suffers from intrinsic non-linearities that require adequate preconditioning. In this study, we investigate a frequency-domain double-beamforming full waveform inversion method (DBF-FWI), applied to acoustic problems. The objective is to take benefit of the array information through a double-beamforming to project the standard point-to-point data to the angles domains on the source and receiver arrays. The DBF procedure allows to improve significantly the signal-to-noise ratio of the DBF data, and the projection in the angles domains allows to select arrivals of interest without any time-domain mute. We illustrate on a synthetic example how the combination of DBF and FWI allows to define a new hierarchical level based on angles, which control image resolution, and that can be used to precondition and improve the robustness of FWI.

INTRODUCTION

Quantitative seismic imaging of subsurface parameters is one of the main challenge for oil and gas reservoir characterization. Full-waveform inversion (FWI) allows to derive high-resolution quantitative models of the subsurface through the exploitation of the full information content of the data (Tarantola, 1984). When applied in the frequency domain, computationally efficient FWI algorithms can be designed by limiting the inversion to a few judiciously chosen discrete frequencies (Sirgue and Pratt, 2004). However, FWI remains an ill-posed and highly non-linear inverse problem that is sensitive to noise, inaccuracies of the starting model, lack of low frequencies in conventional seismic data and definition of multiparameter classes. In order to limit these effects, several hierarchical approaches have been proposed to mitigate the non-linearities of the inverse problem : low to high frequency reconstruction (Bunks et al., 1995; Sirgue and Pratt, 2004), data preconditioning by exponential decay (Brossier et al., 2009a; Brenders et al., 2009), data component and parameter selection (Sears et al., 2008; Brossier et al., 2009b).

Classically, FWI is written as a global summation over the acquisition survey of single source/ single receiver couples. Despite some recent developments tried to use simultaneous source for 3D FWI with source-stacking and phase-encoding strategies (Ben Hadj Ali et al., 2009; Herrmann et al., 2009; Krebs et al., 2009) or plane-wave stacking Vigh and Starr (2008), these approaches do not take benefit of array information and try to mimic the results of point-to-point (PTP) data. However, with the common usage of dense seismic acquisition arrays, one can take benefit of array information to process the seismic data. Techniques such as time-delay beamforming (slant-stack) can therefore be used to identify and select arrivals while improv-

ing the signal-to-noise ratio. Under the source/receiver reciprocity principle, the beamforming can be applied in cascade on both the source and the receiver (sub-)arrays, giving a double-beamforming (DBF) technique (see Rost and Thomas, 2002, for a review).

In this study, we combine the DBF and the acoustic frequency-domain FWI methods. The objective is to make use of the benefits of the DBF advantages, namely the improved signal-to-noise ratio and the arrivals identification/ selection. This method provides a new preconditioning strategy for FWI, based on emission and reception angles on the acquisition array, which allows to control image resolution and should improve the robustness of FWI.

THEORY

FWI recalls

FWI is an optimization problem that is generally recast as a local non-linear least-squares problem. In the frequency domain and for a single frequency, the objective function is :

$$\mathcal{L} = \sum_{k=1}^{ns} \frac{1}{2} \Delta \mathbf{d}_k^t \overline{\Delta \mathbf{d}_k}, \quad (1)$$

where $\Delta \mathbf{d}_k = \mathbf{d}_{obs_k} - \mathbf{d}_{calc_k}$ is the data misfit vector: the difference between the observed data vector \mathbf{d}_{obs_k} and the modeled data vector \mathbf{d}_{calc_k} for the source k . Superscripts t and $^-$ denote the transpose and conjugate operators, respectively. The vector \mathbf{d}_{calc_k} is obtained by applying a sampling operator \mathbf{S} to the full incident wavefield \mathbf{u}_k that results from the forward-problem system $\mathbf{A}\mathbf{u}_k = \mathbf{s}_k$. The matrix \mathbf{A} is the forward-problem matrix that discretizes the Helmholtz equation and which depends non-linearly on the wave-speed model \mathbf{m} . The vector \mathbf{s}_k is the source term that represents the acoustic source excitation. ns is the number of source in the acquisition

The gradient of the cost function can be derived from the adjoint-state formalism (Plessix, 2006):

$$\mathcal{L}_{m_i} = - \sum_{k=1}^{ns} \Re \left\{ \mathbf{u}_k^t \frac{\partial \mathbf{A}^t}{\partial m_i} \overline{\lambda_k} \right\}$$

where \Re denotes the real part of a complex number. The gradient can be seen as a product (zero-lag correlation in time), between the incident wavefield $\mathbf{u}_k = \mathbf{A}^{-1} \mathbf{s}_k$ from the source vector \mathbf{s}_k , and the adjoint wavefield $\overline{\lambda_k} = \mathbf{A}^{-1} \mathbf{S}^t \overline{\Delta \mathbf{d}_k}$, using residuals at receiver positions as a composite source. The radiation pattern of the diffraction by the model parameter m_i is denoted by the sparse matrix $\partial \mathbf{A} / \partial m_i$.

The gradient of the misfit function is then used in an optimization algorithm to update the model vector. In this study, a quasi-Newton L-BFGS (Nocedal, 1980) optimization scheme with line search is used, in the FWI algorithm of Brossier (2011)

Double Beamforming Waveform Inversion

Double beamforming

Double beamforming is a two-step procedure applied on the PTP data in the $(t, \mathbf{x}_s, \mathbf{x}_r)$ domain, where t is time, and \mathbf{x}_s and \mathbf{x}_r are the source and receiver positions. Assuming that the velocity profile is known on the arrays, beamforming (slant-stack) is applied in cascade on source and receiver sub-arrays, projecting the data to the $(t, \mathbf{x}_{s0}, \mathbf{x}_{r0}, \phi, \theta)$ domain, where ϕ and θ are the emission and reception angles on the source and receiver sub-arrays, respectively. Note that DBF is applied on sub-arrays whose size should be optimally close to the first Fresnel zone's size (Iturbe et al., 2009). Finally, a data-selection is applied in the $(t, \mathbf{x}_{s0}, \mathbf{x}_{r0}, \phi, \theta)$ domain. In the present study, the selection has been based on energetic considerations on the envelop of the transformed data. This selection step could be however deeply investigated.

After DBF processing and selection, if we consider one DBF of index k related to the pair of angles (ϕ_k, θ_k) for the sub-arrays centers \mathbf{x}_{r0k} and \mathbf{x}_{s0k} , the frequency-domain expression of the DBF data is given by:

$$d_k = \sum_{r=1}^{n_{rec}} \alpha_{kr} \sum_{s=1}^{n_{src}} \beta_{ks} d_{r,s}, \quad (2)$$

where the summations are performed over the limited numbers of receivers n_{rec} and sources n_{src} . The PTP frequency-domain data $d_{r,s}$ is generated by the source s and measured at the receiver r , and α_{kr} and β_{ks} are the receiver and source phase-shifts associated to the DBF k .

DBF-FWI

The DBF-FWI cost function can be simply written as:

$$\mathcal{C} = \sum_{k=1}^{ndbf} \frac{1}{2} \Delta d_k^t \overline{\Delta d}_k, \quad (3)$$

where the summation is applied on the number of the DBF ($ndbf$) in this case. Each beamforming k is calculated from the source and receiver sub-array centers $(\mathbf{x}_{s0k}, \mathbf{x}_{r0k})$ and the number of sub-array elements n_{src} and n_{rec} , through the associated phase shifts (α_k, β_k) .

After some substitutions, we end up with the expression of the DBF-FWI gradient :

$$\begin{aligned} \mathcal{G}_{m_i} &= - \sum_{k=1}^{ndbf} \Re \left\{ \sum_{r=1}^{n_{rec}} \alpha_{kr} \sum_{s=1}^{n_{src}} \beta_{ks} \mathbf{s}_s^t \mathbf{A}^{-1} \frac{\partial \mathbf{A}^t}{\partial m_i} \mathbf{A}^{-1} \mathbf{S}^t \mathbf{s}_r \overline{\Delta d}_k \right\} \\ &= - \sum_{k=1}^{ndbf} \Re \left\{ \sum_{s=1}^{n_{src}} \beta_{ks} \mathbf{s}_s^t \mathbf{A}^{-1} \frac{\partial \mathbf{A}^t}{\partial m_i} \sum_{r=1}^{n_{rec}} \mathbf{A}^{-1} \mathbf{S}^t \mathbf{s}_r \alpha_{kr} \overline{\Delta d}_k \right\} \end{aligned}$$

where \mathbf{s}_s and \mathbf{s}_r are the impulse source terms located at the source s and receiver r , respectively. As for the classical gradient of Eq. (2), this gradient expression shows the product of three terms:

1. the wavefield $\sum_{s=1}^{n_{src}} \beta_{ks} \mathbf{s}_s^t \mathbf{A}^{-1}$, which represents the incident wavefield from a plane wave located on the source array at position \mathbf{x}_{s0k} with an incident angle θ_{sk} (computed from phase-shifts β_{ks}).

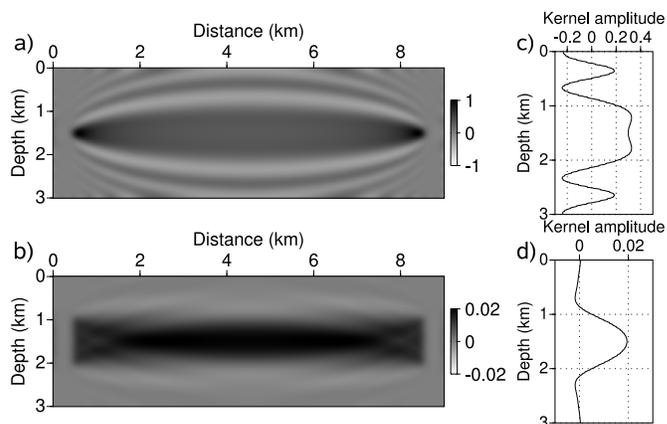


Figure 1: wave-speed sensitivity kernels for (a) PTP acquisition and (b) DBF. Panels (c) and (d) are cross-sections for the kernels at position $x = 4,500$ m for (a) and (b), respectively.

2. the sparse matrix $\partial \mathbf{A}^t / \partial m_i$, which represents the diffraction sensitivity of parameter m_i .
3. the wavefield $\sum_{r=1}^{n_{rec}} \mathbf{A}^{-1} \mathbf{s}_r \alpha_{kr} \overline{\Delta d}_k$, which represents the adjoint wavefield from a plane wave located on the receiver array at position \mathbf{x}_{r0k} , with an incident angle θ_{rk} (computed with phase-shifts α_{kr}), and with an amplitude term $\overline{\Delta d}_k$. This term represents the back-propagation of the residual, as for classical FWI.

SENSITIVITY KERNELS

We first focus on the effects of DBF on sensitivity kernels. Figure 1 illustrates the real part of the PTP and DBF kernels in a homogeneous infinite wave-speed model (1500 m/s) at the frequency of 6.8 Hz. Source and receiver are located at positions $(x=500$ m, $z=1500$ m) and $(x=8500$ m, $z=1500$ m), respectively. The classical PTP kernel is the classical wavepath, which exhibits the first Fresnel zone and also secondary fringes associated with the interference pattern in the monochromatic approach (Woodwards, 1992). The DBF kernel is built by considering 0° incident plane waves on the source and receiver vertical arrays (plane wave propagating perpendicular to survey). Beamforming is applied to 11 elements on each sub-array, which corresponds to 1000 m of antennae. The DBF kernel clearly shows a single sensitivity zone that cannot be related to the first Fresnel zone. This sensitivity zone is built by the constructive summation of each PTP phase-shifted kernel of the source and receiver sub-arrays, and it can be related to plane-wave Fraunhofer diffraction. The secondary fringes of the PTP kernels are destructively summed during the DBF summation, which naturally means that the sensitivity kernel is weaker in the zones where the finite-width plane wave does not propagate. In summary, the DBF-FWI kernel highlights the spatial focusing of kernel sensitivity only where the finite-width plane wave travels, in contrary to the classical FWI kernel, which spreads the sensitivity over a wider area that is associated with the diffraction-based point-spread function of a point-like source-receiver pair.

Double Beamforming Waveform Inversion

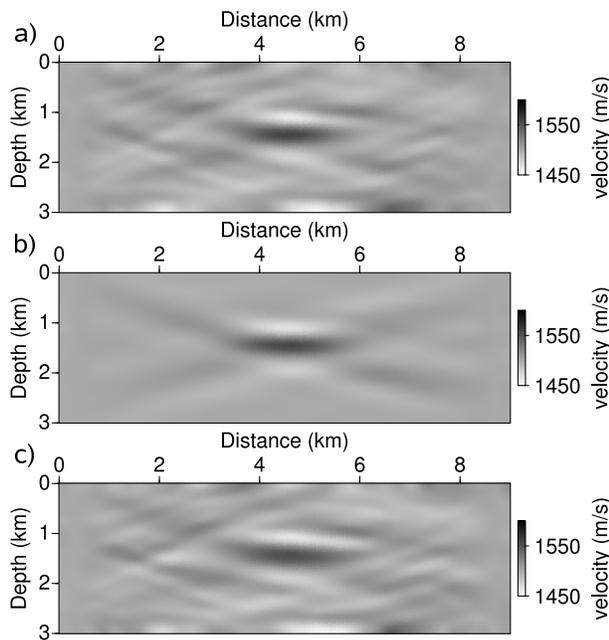


Figure 2: Sound-speed reconstructed model showing the (a) L_2 norm PTP-FWI, (b) DBF-FWI and (c) L_1 norm PTP-FWI for noisy data (SNR=1).

SNR IMPROVEMENT

A first imaging application focuses on a canonical test (single circular anomaly ($V_P = 1750$ m/s) in a homogeneous background ($V_P = 1500$ m/s)) to illustrate how DBF allows to improve the SNR of the reconstructed image tanks to the array gain applied on the data. 29 sources are located on the left side, and 29 receivers on the right side of the model. A mono-frequency of 6.8 Hz is used for inversion. A random noise that affects both the phase and amplitude is added to the PTP data with a SNR of 1.0. This means that the noise has the same power as the useful signal. Figure 2 shows the (a) PTP-FWI and (b) DBF-FWI results in the presence of this noise. DBF has been applied on 11 elements for each sub-array (1000 m). As expected, the results clearly show that the noise introduced into the PTP data strongly affects the model reconstruction in case of L_2 norm minimization. However, the DBF processing mitigates the influence of this noise by the destructive stack of incoherent noise. Using 11-element source-receiver subarrays, an array gain of $20\log_{10}(11) = 20.8$ dB is added to each DBF datum. As a result, the reconstructed wave-speed model is only weakly affected by the noise, providing promising perspectives for real data applications, particularly at low frequencies. Surprisingly, even the L_1 norm (Figure 2-(c)) result appears to be strongly affected by this extreme case of very low SNR.

REALISTIC APPLICATION : OFFSHORE VALHALL

A imaging application focuses on the acoustic synthetic Valhall model (Figure 3-(a)). An acquisition system composed of 315 sources and 315 receivers, each 50 meters, mimics an Ocean Bottom Cable survey. The free-surface is considered in

this study meaning that surface-multiple are present in the data. The Figure 4-(a) shows an example of a shot-gather for a source located at position $x=8100$ m. Inversion is performed for 5 frequencies from 2 Hz to 7 Hz, starting from a smooth starting model (Figure 3-(b)). DBF processing is performed for sub-arrays composed of 21 elements on sources and receivers (1000 meters antennae), and for one central element each 10 elements of the acquisition geometry. Note that DBF could be applied on a finer sub-sample of the survey to improve quality of the results, because this under-sampling of the data leads to a dramatic reduction of the data volume used in FWI.

In order to illustrate how the angles selection on sub-arrays in DBF controls the arrivals taken into account, DBF-FWI is performed in three steps for the angles ranges : $[\pm 60^\circ, \pm 30^\circ]$, $[\pm 60^\circ, \pm 15^\circ]$ and $[\pm 60^\circ, 0^\circ]$. As illustrated on Figure 4, this angle selection allows to choose the arrivals involved in the FWI process without any time-domain windowing despite the frequency-domain formulation. The first DBF range ($[\pm 60^\circ, \pm 30^\circ]$) allows to select the first arrival and very wide angle refractions. The range $[\pm 30^\circ, \pm 15^\circ]$ allows to add pre-critical reflections associated to intermediate diffraction angles, while the final $[\pm 15^\circ, 0^\circ]$ range concerns the short offset reflections associated to small diffraction angles. The resolution power associated to these angles ranges is illustrated in the reconstructed models (Figure 3). We can see, for the same frequency range considered, how the selection on angles on sub-arrays can be linked to the diffraction angles and therefore controls the resolution of the imaged structures. In the framework of FWI where non-linearities should be mitigated with hierarchical procedure, this DBF approach provides an additional multi-scale reconstruction based on angles, which can act as a preconditioner for FWI.

CONCLUSION

This study presents a combined DBF and frequency-domain FWI scheme for subsurface quantitative imaging. The main advantages of this DBF-FWI scheme are linked to the improved SNR of the data thanks to the double summation of the DBF process, which can be particularly useful for low frequency inversion. Under the assumption that the wave-speed is known on the arrays, the method allows also to select the arrivals based on angles that defines a new hierarchical level in order to precondition FWI to mitigate the non-linearities effects, and that could be useful for target-oriented imaging to focus on arrivals of interest. This selection is done without any time-windowing, a great advantage for frequency-domain formulations.

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Double Beamforming Waveform Inversion

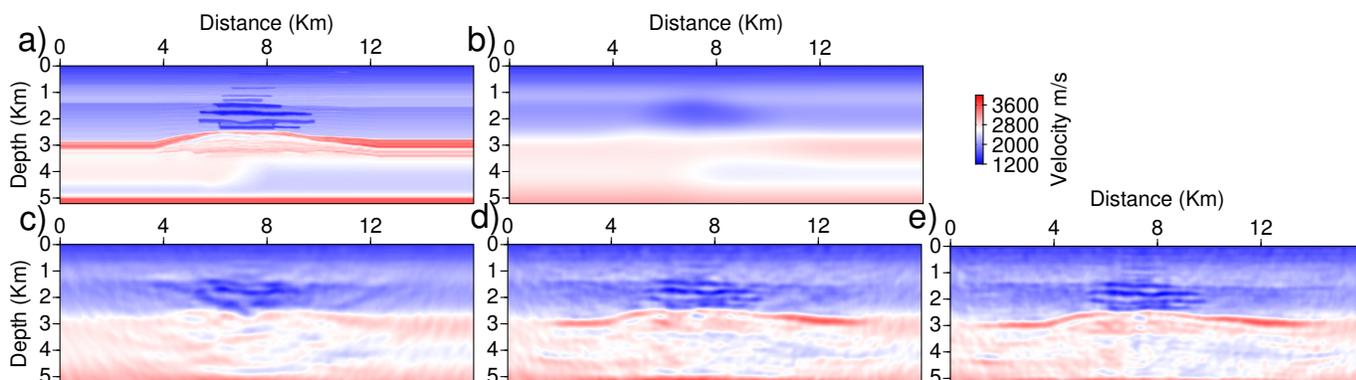


Figure 3: (a) True and (b) initial models of the Valhall application. Reconstructed model after inversion of angle ranges (c) $[\pm 60^\circ, \pm 30^\circ]$, (d) $[\pm 60^\circ, \pm 15^\circ]$ and (e) $[\pm 60^\circ, 0^\circ]$.

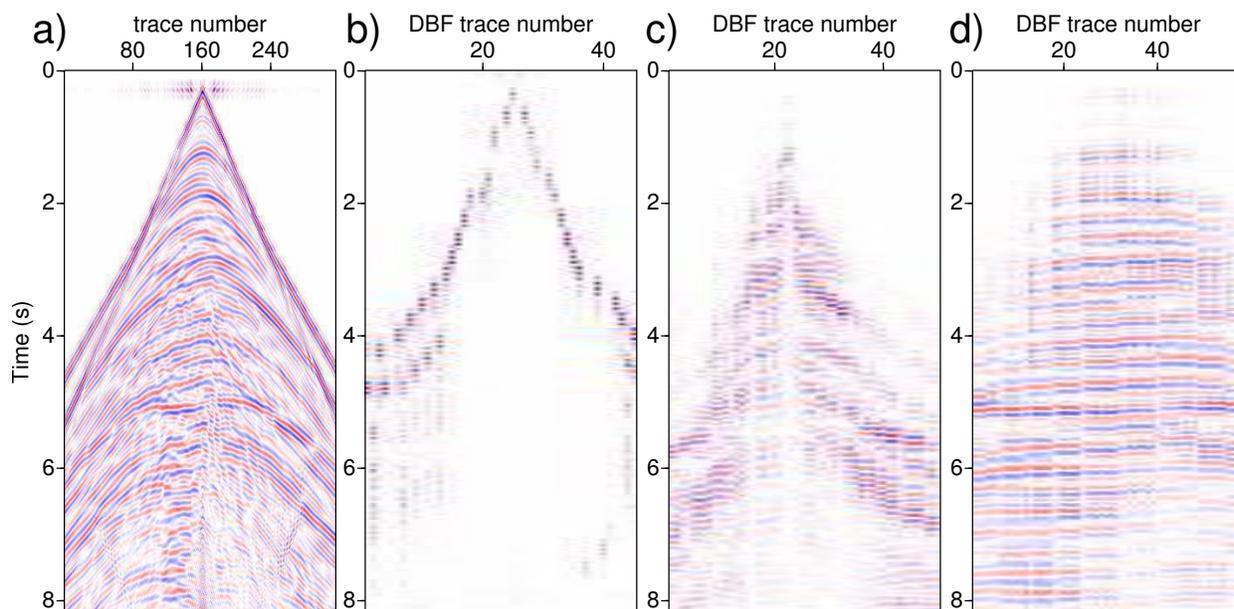


Figure 4: (a) Initial shot gather (source at $x=8100\text{m}$), associated DBF data for the angle ranges (b) $[\pm 60^\circ, \pm 30^\circ]$, (c) $[\pm 30^\circ, \pm 15^\circ]$ and (d) $[\pm 15^\circ, 0^\circ]$. All data are plotted with AGC.

EDITED REFERENCES

Note: This reference list is a copy-edited version of the reference list submitted by the author. Reference lists for the 2011 SEG Technical Program Expanded Abstracts have been copy edited so that references provided with the online metadata for each paper will achieve a high degree of linking to cited sources that appear on the Web.

REFERENCES

- Ben Hadj Ali, H., S. Operto, and J. Virieux, 2009, Three-dimensional frequency-domain full waveform inversion with phase encoding: 79th Annual International Meeting, SEG, Expanded Abstracts, 2288–2292.
- Brenders, A., R. Pratt, and S. Charles, 2009, Waveform tomography of 2-D seismic data in the Canadian foothills — Data preconditioning by exponential time-damping: 71th Conference & Exhibition, EAGE, Extended Abstracts, U041.
- Brossier, R., 2011, Two-dimensional frequency-domain visco-elastic full waveform inversion: Parallel algorithms, optimization and performance: *Computers & Geosciences*, **37**, 444–455, [doi:10.1016/j.cageo.2010.09.013](https://doi.org/10.1016/j.cageo.2010.09.013).
- Brossier, R., S. Operto, and J. Virieux, 2009a, Seismic imaging of complex onshore structures by 2D elastic frequency-domain full-waveform inversion: *Geophysics*, **74**, no. 6, WCC105–WCC118, [doi:10.1190/1.3215771](https://doi.org/10.1190/1.3215771).
- Brossier, R., S. Operto, and J. Virieux, 2009b, Two-dimensional seismic imaging of the Valhall model from synthetic OBC data by frequency-domain elastic full waveform inversion: 79th Annual International Meeting, SEG, Expanded Abstracts, 2293–2297.
- Bunks, C., F. M. Salek, S. Zaleski, and G. Chavent, 1995, Multiscale seismic waveform inversion: *Geophysics*, **60**, 1457–1473, [doi:10.1190/1.1443880](https://doi.org/10.1190/1.1443880).
- Herrmann, F. J., Y. A. Erlangga, and T. T. Y. Lin, 2009, Compressive simultaneous full-waveform simulation: *Geophysics*, **74**, no. 4, A35–A40, [doi:10.1190/1.3115122](https://doi.org/10.1190/1.3115122).
- Iturbe, I., P. Roux, J. Virieux, and B. Nicolas, 2009, Travel-time sensitivity kernels vs. diffraction pattern obtained through double beamforming in shallow water: *The Journal of the Acoustical Society of America*, **126**, 713–720, [doi:10.1121/1.3158922](https://doi.org/10.1121/1.3158922).
- Krebs, J., J. Anderson, D. Hinkley, R. Neelamani, S. Lee, A. Baumstein, and M. D. Lacasse, 2009, Fast full-wavefield seismic inversion using encoded sources: *Geophysics*, **74**, no. 6, WCC177–WCC188, [doi:10.1190/1.3230502](https://doi.org/10.1190/1.3230502).
- Nocedal, J., 1980, Updating Quasi-Newton Matrices With Limited Storage: *Mathematics of Computation*, **35**, no. 151, 773–782, [doi:10.1090/S0025-5718-1980-0572855-7](https://doi.org/10.1090/S0025-5718-1980-0572855-7).
- Plessix, R.-E., 2006, A review of the adjoint-state method for computing the gradient of a functional with geophysical applications: *Geophysical Journal International*, **167**, 495–503, [doi:10.1111/j.1365-246X.2006.02978.x](https://doi.org/10.1111/j.1365-246X.2006.02978.x).
- Rost, S., and C. Thomas, 2002, Array seismology: Methods and applications: *Reviews of Geophysics*, **40**, no. 3, 1008–1035, [doi:10.1029/2000RG000100](https://doi.org/10.1029/2000RG000100).
- Sears, T., S. Singh, and P. Barton, 2008, Elastic full waveform inversion of multi-component OBC seismic data: *Geophysical Prospecting*, **56**, no. 6, 843–862, [doi:10.1111/j.1365-2478.2008.00692.x](https://doi.org/10.1111/j.1365-2478.2008.00692.x).
- Sirgue, L., and R. G. Pratt, 2004, Efficient waveform inversion and imaging: a strategy for selecting temporal frequencies: *Geophysics*, **69**, 231–248, [doi:10.1190/1.1649391](https://doi.org/10.1190/1.1649391).

Tarantola, A., 1984, Inversion of seismic reflection data in the acoustic approximation: *Geophysics*, **49**, 1259–1266, [doi:10.1190/1.1441754](https://doi.org/10.1190/1.1441754).

Vigh, D., and E. W. Starr, 2008, 3D prestack plane-wave, full-waveform inversion: *Geophysics*, **73**, no. 5, VE135–VE144, [doi:10.1190/1.2952623](https://doi.org/10.1190/1.2952623).

Woodward, M. J., 1992, Wave-equation tomography: *Geophysics*, **57**, 15–26, [doi:10.1190/1.1443179](https://doi.org/10.1190/1.1443179).